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The molecular biogeochemistry of manganese(II) oxidation

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Abstract

Micro-organisms capable of oxidizing the redox-active transition metal manganese play an important role in the biogeochemical cycle of manganese. In the present mini-review, we focus specifically on Mn(II)oxidizing bacteria. The mechanisms by which bacteria oxidize Mn(II) include a two-electron oxidation reaction catalysed by a novel multicopper oxidase that produces Mn(IV) oxides as the primary product. Bacteria also produce organic ligands, such as siderophores, that bind to and stabilize Mn(III). The realization that this stabilized Mn(III) is present in many environments and can affect the redox cycles of other elements such as sulfur has made it clear that manganese and the bacteria that oxidize it profoundly affect the Earth's biogeochemistry.

Geochemical cycling of manganese

The carbon and nitrogen cycles may get most of the attention in high school science class, but the manganese cycle also has a tremendous impact on the Earth's geochemistry. Manganese is essential for life as a cofactor in enzymes such as Photosystem II and manganese superoxide dismutase and is one of the most abundant redox-active transition metals in the Earth's crust [1]. In Nature, manganese is commonly found in three oxidation states: Mn(II), Mn(III) and Mn(IV) [2]. The reduced form, Mn(II), is generally soluble and stable in the absence of oxygen. Mn(IV), the most oxidized form, is insoluble, forming oxides which are strong oxidants, capable of oxidizing inorganic and organic compounds, for example, Fe(II) and UO₂ [3]. The intermediate oxidation state, Mn(III), is unstable as an ion under normal environmental conditions unless complexed with organic or inorganic ligands (see below) [1]. Mn(III) also occurs as insoluble Mn(III) oxy(hydrox)oxide phases or in mixed Mn(III,IV) oxides. Mn(III,IV) oxides also sorb metals and other compounds from the environment; as a result, Mn(II) oxidation can control the distribution of many other elements (e.g. copper, cobalt, nickel, lead, iron, radium, uranium and rare earth elements) [4,5].

A one-electron transfer for the oxidation of the soluble Mn(II) species $Mn(H_2O)_6^{2+}$ by O_2 is thermodynamically unfavourable at pH < 9, but a two-electron oxidation is favourable at pH > 3 [6]. Complexed forms of Mn(II) can oxidize at neutral pH, but abiotic Mn(II) oxidation at

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circumneutral pH is quite slow under normal environmental conditions; bacteria and fungi can greatly increase the rate of this reaction by up to 5 orders of magnitude [1]. Thus it is thought that the bulk of environmental Mn(II) oxidation is carried out by micro-organisms. In the present mini-review, we focus specifically on Mn(II)-oxidizing bacteria.

Mn(II)-oxidizing bacteria are widespread and found in diverse environments such as soil, natural waters and sediments. These organisms are phylogenetically diverse, with representatives in the Firmicutes, Actinobacteria and the Alpha-, Beta- and Gamma-proteobacteria [7]. The physiological role Mn(II) oxidation plays in these species is unknown. Because Mn(IV) formation is thermodynamically favourable, the bacteria could derive energy from the reaction; however, this has not been shown conclusively for any organism. The oxidation of Mn(II) may also help to protect the cells from ROS (reactive oxygen species) or other free radicals [8]. Alternatively, the Mn(IV) oxides formed in this reaction are themselves highly reactive and may be used to oxidize refractory organic material that can then be utilized by the micro-organism as a carbon source. Other possible functions for the oxides are as terminal electron acceptors, manganese storage or, since the solids coat the cell, protection from environmental hazards such as UV radiation, predation or phage infection [7,9].

Mn(III,IV) oxide deposits can be found in many environments [1,10]. These include ferromanganese nodules in the deep sea, lakes and soils and terrestrial ferromanganese crusts, also called rock varnish. Mn(IV) oxides are also found in ore deposits and metalliferous sediments associated with spreading centres and as ferromanganese crusts on seamounts on the ocean floor. Given that the oxides reflect the redox conditions of the local environment, they are unsurprisingly

Key words: manganese(II), manganese(III), manganese oxidation, manganese oxide, multicopper oxidase.

Abbreviations used: MCO, multicopper oxidase.

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often found associated with oxic–anoxic interfaces in sediments and water. The extent to which microbial activity is responsible for these formations is unclear. However, the ubiquity of Mn(II)-oxidizing organisms and their presence in environments where Mn(IV) oxides are found suggest that their role is substantial.

Molecular biochemistry of Mn(II) oxidation

Despite the diversity of organisms capable of the reaction, there are common themes in bacterial Mn(II) oxidation [9]. The product is primarily Mn(IV)O2, with oxidation occurring through two sequential one-electron steps $[Mn(II) \rightarrow Mn(III) \rightarrow Mn(IV)]$ that require O₂ [11,12] (Figure 1). Organisms commonly initiate Mn(II) oxidation at the onset of stationary phase and deposit the oxides on their outer surface [1,9]. However, some variations on these common themes occur. Both fungal Mn(II)-oxidizers and the alphaproteobacterium Erythrobacter sp. SD-21 produce Mn(III), which is stabilized by biogenic ligands [13-16]. Another less direct mechanism of Mn(II) oxidation is through the action of enzymatically produced superoxide, as has been observed with the alphaproteobacterium Roseobacter sp. AzwK-3b [17-19]. The iron-chelating siderophores produced by bacteria under iron-limiting conditions are also able to bind with high affinity to manganese; this interaction itself can promote Mn(II) oxidation [20-23].

The enzymes responsible for bacterial Mn(II) oxidation have been identified in several species and fall into two general categories. The alphaproteobacteria *Aurantimonas manganoxydans* SI85-9A1 and *Erythrobacter* sp. SD-21 both employ calcium-binding haem peroxidases named MopA to oxidize Mn(II) [16]. Many other species possess MCO (multicopper oxidase) Mn(II) oxidase enzymes (Figure 2). These include MofA in Leptothrix sp. and the twin Mn(II) oxidase MCOs of Pseudomonas putida GB-1, MnxG and McoA ([24], and Geszvain, K., McCarthy, J.K. and Tebo, B.M., unpublished work). MCO enzymes catalyse the oneelectron transfer from substrates as diverse as Fe(II) and lignin to ultimately reduce O₂ to H₂O [26]. MCOs incorporate multiple copper ligands, including a type I copper-binding site and a trinuclear centre made up of two type II and one type III site. One electron is removed from the substrate at the type I copper then shuttled to the trinuclear centre where it is ultimately added to O2 to form H2O. As the diversity of substrates would suggest, this family of enzymes is very diverse; identification of family members is through conservation of characteristic copper-binding sites, identified both at the sequence level and by their spectroscopic characteristics [26,27].

The Mn(II) oxidase from *Bacillus* sp., MnxG, is the best characterized oxidase enzyme to date. MnxG was first implicated as the *Bacillus* Mn(II) oxidase via transposon mutagenesis. Several transposon insertions were identified within an operon of eight genes (*mnxA-mnxG*) that resulted in mutant strains that failed to oxidize Mn(II) [28,29]. The last gene in the operon, *mnxG*, encodes an MCO, based on the presence of five putative copper-binding domains. The previous observation that the addition of copper stimulated Mn(II) oxidation supported the idea that an MCO was the Mn(II) oxidase in this organism [29]. Subsequent partial purification of the Mn(II) oxidase activity from *Bacillus* sp. PL-12 exosporium followed by mass spectrophotometry identified MnxG in the active fraction, supporting the identification of this MCO as the Mn(II) oxidase [30,31].

MCO enzymes catalyse the one-electron oxidation of their substrates, whereas Mn(II) oxidation appears to require two electron transfers, making the mechanism of Mn(II) oxidation a subject of intense study. This mechanism has been investigated using exosporium prepared from Bacillus sp. SG-1 [30-32]. Exosporium preparations are capable of oxidizing Mn(II) to Mn(IV); in the presence of pyrophosphate, a stable Mn(III)-pyrophosphate complex forms that can be monitored by measuring the absorbance at 258 nm [32]. From these experiments, it was concluded that Mn(II) oxidation in fact proceeds through an Mn(III) intermediate [32] and that oxygen was required for each step [12]. Both the Mn(II) to Mn(III) and Mn(III) to Mn(IV) reactions were absent from exosporium prepared from a strain in which mnxG was disrupted by a transposon insertion [30]. Therefore MnxG carries out each electron transfer, differentiating Mn(II) oxidase MCOs from the typical MCO which performs just one electron transfer.

The large size and domain organization of MnxG makes it somewhat similar to human caeruloplasmin, a ferroxidase [29,30]. Previous crystallographic and biochemical studies of caeruloplasmin make it possible to propose a model of the mechanism of Mn(II) oxidation by MnxG [8] (Figure 3). In caeruloplasmin, Fe(II) binds to a site adjacent to the type 1 copper; after electron transfer, the product Fe(III) moves several angstroms away to a 'holding site'. MnxG

Figure 2 | Copper-binding motifs from Mn(II) oxidases and other MCOs

Residues implicated in copper binding are highlighted. Hcer, human caeruloplasmin; Lacc, laccase.

Region A										Region B													
BaMnxG	527	м	н	1	H	ł	F V			BaMnxG	572	F	F	н	D	н		Bacillus sp. SG-1					
PpMnxG	942	Q	н	1	H	1	LP			PpMnxG	1045	F	т	н	D	н		P. putida GB-1					
McoA	343	т	н	E	H		N	G		McoA	435	w	F	н	D	н		P. putida GB-1					
MofA	304	Т	н	L	H	•	G	G		MofA	384	w	Υ	н	D	н		L. discophora SS-1					
MoxA	129	Т	н	w	H	•	G	Q		MoxA	170	М	Υ	н	Р	н	Pedomicrobium			n AC	ACM3067		
Hcer	119	F	н	F	H	•	G	L		Hcer	178	1	Υ	н	s	н	1	Human					
Lacc	78	V	н	w	H	ł.	G	L		Lacc	121	w	Υ	н	s	н	1	Fungi/plants					
Region C										Region D													
BaMnxG	281	н	v	F	н	Υ	н	v	н	BaMnxG	334	н	С	н	L	Υ	Ρ	н	F	G	Т	G	м
PpMnxG	477	н	1	F	н	L	н	G	н	PpMnxG	553	н	С	н	F	Υ	Ρ	н	F	А	Q	G	м
McoA	816	н	Р	v	н	v	н	F	Е	McoA	874	н	С	н	Ν	т	Q	н	Е	D	s	s	М
MofA	1174	н	Ρ	v	н	F	н	L	L	MofA	1279	н	С	н	1	L	G	н	Е	Е	Ν	D	F
MoxA	265	н	Р	Т	н	м	н	G	Υ	MoxA	318	н	С	н	к	s	н	н	т	м	Ν	Α	м
Hcer	994	н	т	v	н	F	н	G	н	Hcer	1039	н	С	н	v	т	D	н	I.	н	А	G	м
Lacc	508	н	Ρ	I.	н	к	н	G	Ν	Lacc	585	н	С	н	I	А	s	н	Q	м	G	G	М

may similarly oxidize Mn(II) to Mn(III). The presence of a solvent-accessible Mn(III) holding site on MnxG could account for the complexation of Mn(III) by pyrophosphate. Since mononuclear Mn(IV) is a powerful oxidant, MnxG probably forms a polynuclear Mn(IV) oxide. Thus it has been proposed that MnxG has a binuclear Mn(III) site adjacent to the holding site, with a third Mn(III) ion serving as an electron shuttle between the binuclear site and the type I copper [12]. Future work with purified MnxG will make it possible to investigate in detail the molecular mechanism of Mn(II) oxidation.

Environmental geochemistry of soluble Mn(III)

Much work on the environmental roles of manganese species has been carried out in the Black Sea, a useful environmental laboratory due to its permanent halocline, a broad suboxic layer characterized by O_2 levels below 5 μ M and undetectable H_2S and an anoxic H_2S -rich layer below ~150 m. Therefore the Black Sea can be used as a model oxic-suboxicanoxic transition zone. The suboxic zone is thought to be maintained by microbially driven, O₂-dependent Mn(II) oxidation depleting the last traces of O2, whereas the Mn(IV) oxides serve to cap the vertical flux of H₂S. However, the O₂ consumption to form manganese oxides is not sufficient to account for the H₂S depletion. Originally, it was proposed that Mn(II) oxidation was coupled to nitrate reduction, but no system has yet been identified in which nitrate serves as the electron acceptor in Mn(II) oxidation; in all cases studied so far, oxidation has required O_2 .

Mn(III) historically has not been considered to be important in the environment because it was thought to be too

reactive to exist in natural environments. Furthermore, manganese chemistry has been operationally defined, with Mn(II) being the species that passes through 0.4 μ m filters (i.e. soluble), whereas Mn(III,IV) oxides would be trapped on the filters. However, Mn(III) stabilized by complexation to organic or inorganic ligands can be soluble, pass through filters and be measured as Mn(II). Work in the laboratory suggested that Mn(III) complexed with siderophores is stable and soluble, raising the possibility that such soluble Mn(III) is an even more important redox agent in the environment than O₂.

To determine soluble Mn(III) levels in the Black Sea, it was necessary to devise methods to routinely detect it in environmental samples. A sensitive technique to measure Mn(III) in water samples by employing the addition of the siderophore desferrioxamine B to complex Mn(III) and subsequent quantification using cyclic voltammetry was developed [33]. With this method, it was shown that Mn(III) levels in the suboxic zone were in fact quite high (4–5 μ M) and constituted up to 100% of the dissolved manganese (passing through a $0.2 \,\mu m$ filter). The ligand-complexed Mn(III), which could be both a reductant and an oxidant, could therefore be responsible for capping the H₂S flux in the suboxic zone. Subsequent modelling results support the conclusion that the H₂S is consumed by oxidized forms of manganese with no need to propose the presence of alternative electron acceptors such as nitrate [34,35].

Similar work has been carried out with sediments collected along the Lower Saint Lawrence Estuary, using a manganese speciation method based on a metal substitution reaction with soluble porphyrin [11,36]. This work also detected up to $65 \,\mu\text{M}$ soluble Mn(III), therefore the effect of this species on environmental redox cycling may be widespread and

Figure 3 | Proposed mechanism of bacterial Mn(II) oxidation

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profound. The nature of the ligands responsible for stabilizing soluble Mn(III) is unknown. However, in the Black Sea, Mn(II)-oxidizing bacteria have been isolated that also produce siderophores that are capable of binding Mn(III) [37].

Siderophores have been traditionally studied for their role in iron sequestration and uptake. The realization that they also bind manganese, resulting in high concentrations of Mn(III) in the environment, makes it necessary to rethink the role that they play in metal cycling. In many regions of the ocean, levels of available Fe(III) are low and levels of ironcomplexing ligands resembling siderophores are high. Clearly, the cycling of manganese, as is evident from recent studies of manganese oxidation, is mechanistically much more complex than depicted by the paradigm of manganese that passes through a filter is dissolved Mn(II) and that retained by the filter is Mn(III,IV) oxide (Figure 1).

Conclusions

Mn(II)-oxidizing bacteria are widespread and diverse, and the enzymes directly responsible appear to fall in to two general categories: haem peroxidase and MCO. Indirect oxidation of Mn(II) may also occur through the enzymatic production of superoxide. Soluble Mn(III) has been found in a variety of environments and is expected to be present wherever there is a redox gradient. As both an oxidant and a reductant, soluble Mn(III) can profoundly affect redox cycling in complex ways. Furthermore, since iron-chelating siderophores also bind Mn(III), there is competition between Fe(III) and Mn(III) for the same ligands, complicating the effect of soluble Mn(III) on iron biogeochemistry. Mn(III) may also be linked to other biogeochemical cycles, such as nitrogen, sulfur and phosphorus, but the importance of the coupling of manganese cycling with these other elemental cycles is completely unknown. Clearly, study of Mn(II) oxidation at the molecular, biochemical and environmental levels has only just begun.

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